

NOTES ON BASE

This map sheet is one of a series covering the entire surface of Mars at nominal scales of 1:25,000,000 and 1:5,000,000 (Barton, 1972, 1976). The major sources of map data were the Mariner 9 television experiment (Masursky and others, 1970), Mariner 9 photography (1971-1972), and the Mars Global Surveyor (MGS) Mars Orbiter Camera (MOC) (1996-1999).

The figure of Mars used for the compilation of the map projection is an oblique spheroid with a semi-major axis of 2332.4 km and a semi-minor axis of 2332.2 km. The equatorial radius of 2332.4 km and a polar radius of 2332.2 km are used for the figure of Mars which is defined below under the heading "CONTOURS".

PROJECTION
The Mercator projection is used for this sheet, with a scale of 1:5,000,000 at the equator and 1:4,266,000 at lat 30° N. Longitudinal scale increases to the north to coincide with the map of the International Astronomical Union (IAU), 1971. Latitude is geographic (i.e., true) and not geocentric (IAU, 1973).

CONTROL
Planimetric control is provided by photogrammetric triangulation using Mariner 9 pictures (Davies, 1973; Davies and Arthur, 1973) and the radio-tracked position of the spacecraft. The first meridian passes through the crater Aps 0 (lat 51° S) within the crater area. No simple statement is possible for the precision, but local consistency is ± 0.10 mm.

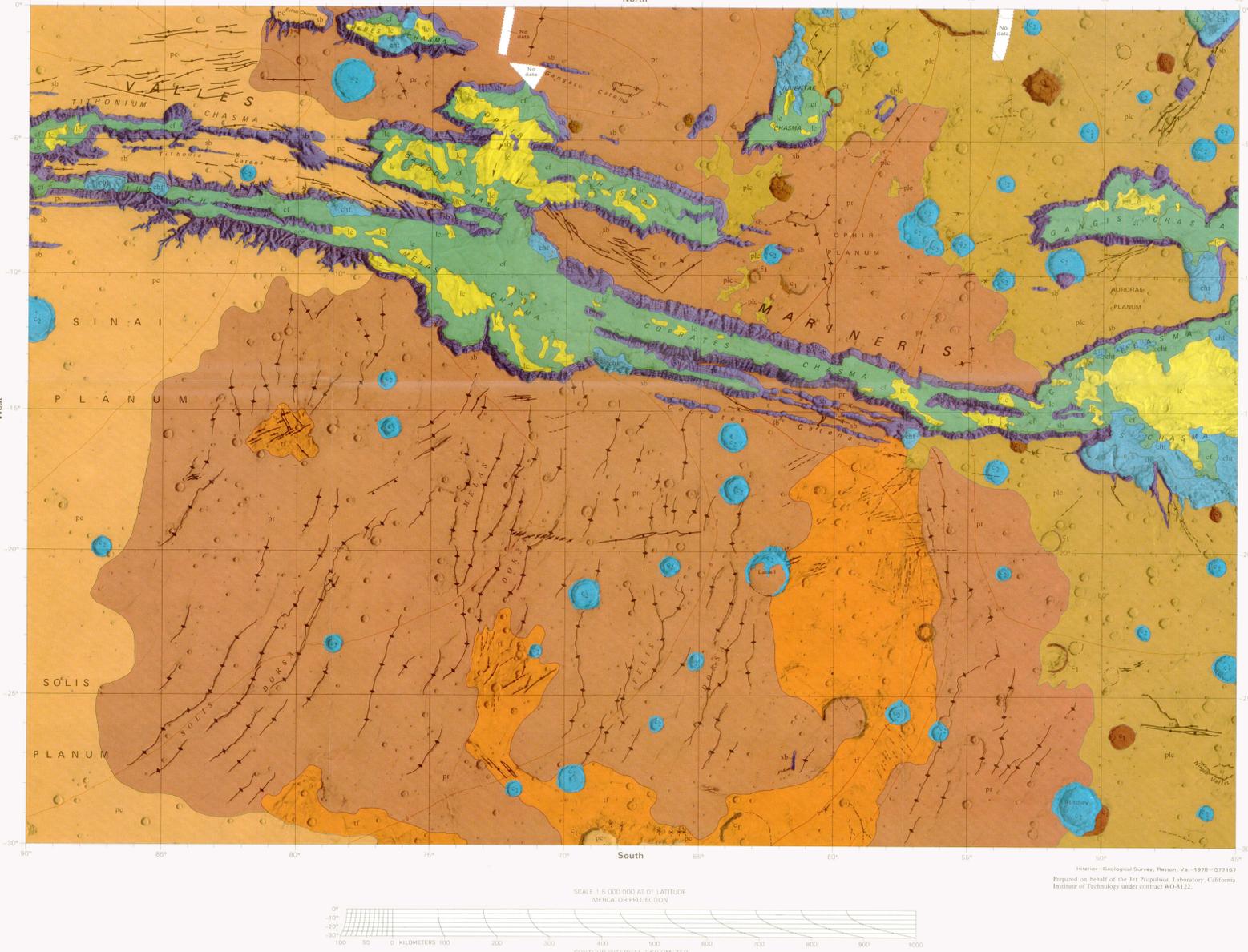
MAPPING TECHNIQUES
A series of mosaics of Mariner 9 photographs of Mars 9 pictures was assembled at 1:5,000,000. Shaded relief was copied from the mosaics and portrayed with uniform illumination with the sun on the west. Many Mariner 9 pictures besides those in the base mosaic were examined to determine the purpose of the shaded relief (1973; Green and others, 1973; Inge and Hindsley, 1976). The shading is not generated and does not interfere with astrophotographic reliability (Inge, 1972).

CONTOURS
Since Mars has no sea and hence no sea level, the datum (the 0 km contour line) for altitudes is defined by a gravity field model (Jordan and Loebel, 1973) combined with a 6.1-millibar atmospheric pressure surface derived from altimetry data (Kjore and others, 1973; Christensen, 1975; Wu, 1975).

The contour lines on most of the Mars maps (Wu, 1975) were computed from altimetry data derived from the Mars Global Surveyor (MGS) Mars Orbiter Camera (MOC) (1996-1999) and measurements made by the Mars Global Surveyor (MGS) Mars Orbiter Camera (MOC) (1996-1999) and measurements made by the Mars Global Surveyor (MGS) Mars Orbiter Camera (MOC) (1996-1999).

All names on this sheet are approved by the International Astronomical Union (IAU), 1971. The names of Mars 9 pictures are given in the margin of this sheet. The names of Mars 9 pictures are given in the margin of this sheet. The names of Mars 9 pictures are given in the margin of this sheet.

REFERENCES
Barton, R. M., 1973. Cartographic products from the Mariner 9 mission. *Jour. Geophys. Research*, v. 78, no. 20, p. 4424-4436.
Christensen, P. R., 1975. Martian topography derived from altimetry data. *Jour. Geophys. Research*, v. 80, no. 20, p. 2909-2911.
Green, H., 1973. Cartographic products from the Mariner 9 mission. *Jour. Geophys. Research*, v. 78, no. 20, p. 4424-4436.



CORRELATION OF MAP UNITS

CHASMA MATERIALS	OTHER MATERIALS	CRATER MATERIALS
ch	pc	ch
ch1	pc1	ch1
ch2	pc2	ch2
ch3	pc3	ch3
ch4	pc4	ch4
ch5	pc5	ch5
ch6	pc6	ch6
ch7	pc7	ch7
ch8	pc8	ch8
ch9	pc9	ch9
ch10	pc10	ch10
ch11	pc11	ch11
ch12	pc12	ch12
ch13	pc13	ch13
ch14	pc14	ch14
ch15	pc15	ch15
ch16	pc16	ch16
ch17	pc17	ch17
ch18	pc18	ch18
ch19	pc19	ch19
ch20	pc20	ch20
ch21	pc21	ch21
ch22	pc22	ch22
ch23	pc23	ch23
ch24	pc24	ch24
ch25	pc25	ch25
ch26	pc26	ch26
ch27	pc27	ch27
ch28	pc28	ch28
ch29	pc29	ch29
ch30	pc30	ch30
ch31	pc31	ch31
ch32	pc32	ch32
ch33	pc33	ch33
ch34	pc34	ch34
ch35	pc35	ch35
ch36	pc36	ch36
ch37	pc37	ch37
ch38	pc38	ch38
ch39	pc39	ch39
ch40	pc40	ch40
ch41	pc41	ch41
ch42	pc42	ch42
ch43	pc43	ch43
ch44	pc44	ch44
ch45	pc45	ch45
ch46	pc46	ch46
ch47	pc47	ch47
ch48	pc48	ch48
ch49	pc49	ch49
ch50	pc50	ch50
ch51	pc51	ch51
ch52	pc52	ch52
ch53	pc53	ch53
ch54	pc54	ch54
ch55	pc55	ch55
ch56	pc56	ch56
ch57	pc57	ch57
ch58	pc58	ch58
ch59	pc59	ch59
ch60	pc60	ch60
ch61	pc61	ch61
ch62	pc62	ch62
ch63	pc63	ch63
ch64	pc64	ch64
ch65	pc65	ch65
ch66	pc66	ch66
ch67	pc67	ch67
ch68	pc68	ch68
ch69	pc69	ch69
ch70	pc70	ch70
ch71	pc71	ch71
ch72	pc72	ch72
ch73	pc73	ch73
ch74	pc74	ch74
ch75	pc75	ch75
ch76	pc76	ch76
ch77	pc77	ch77
ch78	pc78	ch78
ch79	pc79	ch79
ch80	pc80	ch80
ch81	pc81	ch81
ch82	pc82	ch82
ch83	pc83	ch83
ch84	pc84	ch84
ch85	pc85	ch85
ch86	pc86	ch86
ch87	pc87	ch87
ch88	pc88	ch88
ch89	pc89	ch89
ch90	pc90	ch90
ch91	pc91	ch91
ch92	pc92	ch92
ch93	pc93	ch93
ch94	pc94	ch94
ch95	pc95	ch95
ch96	pc96	ch96
ch97	pc97	ch97
ch98	pc98	ch98
ch99	pc99	ch99
ch100	pc100	ch100

DESCRIPTION OF MAP UNITS

CHASMA MATERIALS
SLOPE AND BEDROCK MATERIAL - Composite unit that occurs on relatively steep slopes along walls of chasmas, chain craters, and collapse features. Layering in form of dark band along hallway up chasma walls seen on south side of Coprates. Slopes average 20° to 32° over most of the chasma walls. Walls of Vallis Marineris show signs of erosion patterns that include steep headward downward bifurcating ridges and gullies, and large alcoves below which lie great masses of jumbled debris. Gullies generally produced by regional structure pattern present along with large inclined bifurcating sidewall gullies. Upper slopes commonly marked by vertical ribs and grooves. Large flow lobes common at base of scarp. Representative locality in northern part of Coprates Chasma at long 89° 44', lat 23° 12' (DAS 7042584), where ten or more alternating light and dark layers are exposed mostly by mass wasting or gravitational collapse of over-steepened slopes. Large bifurcating ridges produced by rapid flowage of water-saturated sediments after breaching of isolated lake basins within Vallis Marineris. Lack of benching in chasma subfloors suggests homogeneity in vertical section.

CHASMA FLOOR MATERIAL - Occurs on floor of Vallis Marineris in subdividing stratified mesas. Representative locality in northern part of Coprates Chasma at long 89° 44', lat 23° 12' (DAS 7042584), where ten or more alternating light and dark layers are exposed mostly by mass wasting or gravitational collapse of over-steepened slopes. Large bifurcating ridges produced by rapid flowage of water-saturated sediments after breaching of isolated lake basins within Vallis Marineris. Lack of benching in chasma subfloors suggests homogeneity in vertical section.

CHASMA FLOOR MATERIAL - Occurs on floor of Vallis Marineris in subdividing stratified mesas. Representative locality in northern part of Coprates Chasma at long 89° 44', lat 23° 12' (DAS 7042584), where ten or more alternating light and dark layers are exposed mostly by mass wasting or gravitational collapse of over-steepened slopes. Large bifurcating ridges produced by rapid flowage of water-saturated sediments after breaching of isolated lake basins within Vallis Marineris. Lack of benching in chasma subfloors suggests homogeneity in vertical section.

CHASMA FLOOR MATERIAL - Occurs on floor of Vallis Marineris in subdividing stratified mesas. Representative locality in northern part of Coprates Chasma at long 89° 44', lat 23° 12' (DAS 7042584), where ten or more alternating light and dark layers are exposed mostly by mass wasting or gravitational collapse of over-steepened slopes. Large bifurcating ridges produced by rapid flowage of water-saturated sediments after breaching of isolated lake basins within Vallis Marineris. Lack of benching in chasma subfloors suggests homogeneity in vertical section.

CHASMA FLOOR MATERIAL - Occurs on floor of Vallis Marineris in subdividing stratified mesas. Representative locality in northern part of Coprates Chasma at long 89° 44', lat 23° 12' (DAS 7042584), where ten or more alternating light and dark layers are exposed mostly by mass wasting or gravitational collapse of over-steepened slopes. Large bifurcating ridges produced by rapid flowage of water-saturated sediments after breaching of isolated lake basins within Vallis Marineris. Lack of benching in chasma subfloors suggests homogeneity in vertical section.

CHASMA FLOOR MATERIAL - Occurs on floor of Vallis Marineris in subdividing stratified mesas. Representative locality in northern part of Coprates Chasma at long 89° 44', lat 23° 12' (DAS 7042584), where ten or more alternating light and dark layers are exposed mostly by mass wasting or gravitational collapse of over-steepened slopes. Large bifurcating ridges produced by rapid flowage of water-saturated sediments after breaching of isolated lake basins within Vallis Marineris. Lack of benching in chasma subfloors suggests homogeneity in vertical section.

CHASMA FLOOR MATERIAL - Occurs on floor of Vallis Marineris in subdividing stratified mesas. Representative locality in northern part of Coprates Chasma at long 89° 44', lat 23° 12' (DAS 7042584), where ten or more alternating light and dark layers are exposed mostly by mass wasting or gravitational collapse of over-steepened slopes. Large bifurcating ridges produced by rapid flowage of water-saturated sediments after breaching of isolated lake basins within Vallis Marineris. Lack of benching in chasma subfloors suggests homogeneity in vertical section.

CHASMA FLOOR MATERIAL - Occurs on floor of Vallis Marineris in subdividing stratified mesas. Representative locality in northern part of Coprates Chasma at long 89° 44', lat 23° 12' (DAS 7042584), where ten or more alternating light and dark layers are exposed mostly by mass wasting or gravitational collapse of over-steepened slopes. Large bifurcating ridges produced by rapid flowage of water-saturated sediments after breaching of isolated lake basins within Vallis Marineris. Lack of benching in chasma subfloors suggests homogeneity in vertical section.

CHASMA FLOOR MATERIAL - Occurs on floor of Vallis Marineris in subdividing stratified mesas. Representative locality in northern part of Coprates Chasma at long 89° 44', lat 23° 12' (DAS 7042584), where ten or more alternating light and dark layers are exposed mostly by mass wasting or gravitational collapse of over-steepened slopes. Large bifurcating ridges produced by rapid flowage of water-saturated sediments after breaching of isolated lake basins within Vallis Marineris. Lack of benching in chasma subfloors suggests homogeneity in vertical section.

GEOLOGIC SETTING

The Coprates quadrangle lies to the east of the Tharsis volcanic complex and the Noctis Labyrinthus tectonic complex. To the north are the plains of Lunae Planum. To the east the vast patches of chaotic terrain and the large sinuous channels that trend northward into Chryse Planitia. On the southwest is the Thaumasia trough region of Icthyonically deformed cratered terrain, to the southeast, the Argyre Basin, the best preserved of the ancient martian oceans. To the south are the plains of the Coprates quadrangle. It is dominated by the Vallis Marineris chasma system, which stretches an east-southeast direction for about 2,500 km across the quadrangle. Its maximum width is about 600 km from the north rim of Ophir Chasma to the south rim of Melas Chasma. The terminus of Vallis Marineris at the eastern edge of the map is about 500 km wide in the north-south direction. Individual segments of the chasma system are 100 to 150 km wide and 2 to 4 km deep.

Much of the geology described here is common to other regions of Mars and has been discussed previously not only in summary papers but by other quadrangle workers. Preliminary attempts at regional geology of Mars have been made by McCauley and others (1972) and Carr and others (1973). These units and features common to other parts of Mars are discussed first. The focus of this discussion is on Vallis Marineris and those features that are unique to this quadrangle.

GENERAL STRATIGRAPHY
Moderately cratered plateau material occurs along the eastern edge of the mapped area and marks the beginning of the cratered hemisphere of Mars. This unit bears some resemblance to the lunar basal but is not saturated with large craters, and the intercrater area is smoother and more extensive. Consistent with the appearance of this terrain, the Vallis Marineris chasma system winds markedly, the chasma walls become lower in height, and chaotic terrain is abundant. Sinuous channels are numerous, and the floors of many of the large craters such as that south of Coprates Chasma (IP 8, S17 90° 30' E) have been collapsed. Subsurface sapping and fluvial activity are more widespread in this unit than elsewhere in the quadrangle. Sharp (1973) and Sharp and Main (1975) along with Milton (1974) have suggested mechanisms by which such processes operate on Mars.

The central part of the quadrangle is mostly occupied by ridged plains that are continuous with and similar to the materials of Lunae Planum described by Milton (1974). These plains do not extend much beyond the southern boundary of the map area but patches may be present in the Thaumasia quadrangle. The ridged plains of the Coprates quadrangle and Lunae Planum to the north are not clearly related to large melting events or to craters, as are most of the lunar maria. The ridged plains appear to be volcanic in origin, and the ridges are mostly north-south-southwest and northeast-southeast-southwest. The ridges appear to be volcanic in origin, and the ridges are mostly north-south-southwest and northeast-southeast-southwest. The ridges appear to be volcanic in origin, and the ridges are mostly north-south-southwest and northeast-southeast-southwest.

Several large groups of trough-and-floor material embayed by the ridged plains extend northward into the central part of the quadrangle. They appear to be extensions of the tectonically deformed terrain of the Thaumasia quadrangle described by McGill (1978). Although this occurrence is not obviously stratified, the lobes form on north-south trending ridges that identified as troughed and furrowed hills mapped away from Vallis Marineris. The cratered plains in the southern part of the quadrangle have bright streaks that trend south-southwest. In the same unit in the northern part of the map area, they trend south-southwest toward the lower ends of Vallis Marineris. The dark streaks of the Fella Dorsa region have a general trend that is west-southwest. Many of these markings appear to follow topography in the quadrangle direction, suggesting that slope winds may be a factor in their development.

GEOLOGIC HISTORY
The earliest recognizable geologic events in the Coprates quadrangle are associated with the formation and modification of the cratered plains material on the east side of the map area. This region is part of the more heavily cratered northern hemisphere of Mars consisting of weathered lunar impact craters separated by broad tracts of smooth rolling plains or hilly terrain. The origin of these plains is not understood, although it is probably related to the tectonic evolution of the crust of the order of several kilometers to obliterate almost all traces of the original impact craters. The ridged plains are volcanic in origin, and the ridges are mostly north-south-southwest and northeast-southeast-southwest. The ridges appear to be volcanic in origin, and the ridges are mostly north-south-southwest and northeast-southeast-southwest.

VALLIS MARINERIS
Vallis Marineris has no known terrestrial or planetary counterpart. This structure in the major resembles neither a major rift system nor a valley produced by fluvial downcutting. It is characterized by several long segments such as the Coprates and Coprates Chasma. The Coprates Chasma to the north are more like a normal fault system, but the main part of Vallis Marineris by gaps in the chasma walls. Hebes, Candor, and Ganis Chasmas are separate large depressions that are not integrated into the main system. Sharp (1974), in comparing his Chasma with the Hebes system, noted similarities of form and tectonic pattern. Mason (1977) has shown that the general tectonic pattern is not the present appearance of the Late African Rift is similar to that observed in the Hebes system. The general tectonic pattern is not the present appearance of the Late African Rift is similar to that observed in the Hebes system. The general tectonic pattern is not the present appearance of the Late African Rift is similar to that observed in the Hebes system.

CHASMA MATERIALS
The walls of the chasma that make up the system have been extensively modified by erosion (Sharp, 1973) that has produced extensive lateral and longitudinal enlargement of Vallis Marineris since the system first began to form. The walls of the chasma that make up the system have been extensively modified by erosion (Sharp, 1973) that has produced extensive lateral and longitudinal enlargement of Vallis Marineris since the system first began to form. The walls of the chasma that make up the system have been extensively modified by erosion (Sharp, 1973) that has produced extensive lateral and longitudinal enlargement of Vallis Marineris since the system first began to form.

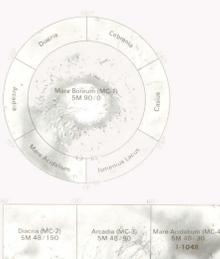
CHASMA MATERIALS
The walls of the chasma that make up the system have been extensively modified by erosion (Sharp, 1973) that has produced extensive lateral and longitudinal enlargement of Vallis Marineris since the system first began to form. The walls of the chasma that make up the system have been extensively modified by erosion (Sharp, 1973) that has produced extensive lateral and longitudinal enlargement of Vallis Marineris since the system first began to form. The walls of the chasma that make up the system have been extensively modified by erosion (Sharp, 1973) that has produced extensive lateral and longitudinal enlargement of Vallis Marineris since the system first began to form.

CHASMA MATERIALS
The walls of the chasma that make up the system have been extensively modified by erosion (Sharp, 1973) that has produced extensive lateral and longitudinal enlargement of Vallis Marineris since the system first began to form. The walls of the chasma that make up the system have been extensively modified by erosion (Sharp, 1973) that has produced extensive lateral and longitudinal enlargement of Vallis Marineris since the system first began to form. The walls of the chasma that make up the system have been extensively modified by erosion (Sharp, 1973) that has produced extensive lateral and longitudinal enlargement of Vallis Marineris since the system first began to form.

CHASMA MATERIALS
The walls of the chasma that make up the system have been extensively modified by erosion (Sharp, 1973) that has produced extensive lateral and longitudinal enlargement of Vallis Marineris since the system first began to form. The walls of the chasma that make up the system have been extensively modified by erosion (Sharp, 1973) that has produced extensive lateral and longitudinal enlargement of Vallis Marineris since the system first began to form. The walls of the chasma that make up the system have been extensively modified by erosion (Sharp, 1973) that has produced extensive lateral and longitudinal enlargement of Vallis Marineris since the system first began to form.

ALBEDO MARKINGS AND CONTOURS

Contour interval 1 kilometer. Surface markings derived from selected Mariner 9 photographs.



INDEX TO MARINER 9 PICTURES

The mosaic used to control the positions of features on this map was made with the Mariner 9 camera pictures outlined above. Unfilled coverage is not available in some craters shown. Pictures other than those shown in the mosaic were used for parallax in the other craters marked. Also shown by solid black rectangles are the high-resolution cameras pictures, identified by italic numbers.

Camera picture	High resolution camera picture
000001	000001
000002	000002
000003	000003
000004	000004
000005	000005
000006	000006
000007	000007
000008	000008
000009	000009
000010	000010
000011	000011
000012	000012
000013	000013
000014	000014
000015	000015
000016	000016
000017	000017
000018	000018
000019	000019
000020	000020
000021	000021
000022	000022
000023	000023
000024	000024
000025	000025
000026	000026
000027	000027
000028	000028
000029	000029
000030	000030
000031	000031
000032	000032
000033	000033
000034	000034
000035	000035
000036	000036
000037	000037
000038	000038
000039	000039
000040	000040
000041	000041
000042	000042
000043	000043
000044	000044
000045	000045
000046	000046
000047	000047
000048	000048
000049	000049
000050	000050
000051	000051
000052	000052
000053	000053
000054	000054
000055	000055
000056	000056
000057	000057
000058	000058
000059	000059
000060	000060
000061	000061
000062	000062
000063	000063
000064	000064
000065	000065
000066	000066
000067	000067
000068	000068
000069	000069
000070	000070
000071	000071
000072	000072
000073	000073
000074	000074
000075	000075
000076	000076
000077	000077
0000	